

# Experimental and numerical modelling of internal solitary waves

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# Overview

- Background & Motivation
- Experimental Method
- Numerical Method
- Some Examples
- A little bit about me & my career

# Solitary Waves



John Scott Russell (1834)



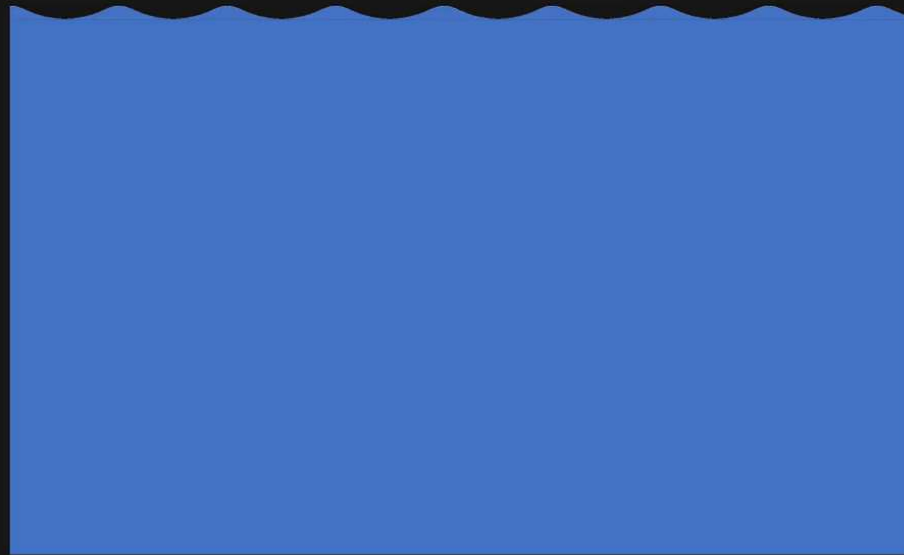
Scott Russell Aqueduct



Scott Russell Aqueduct

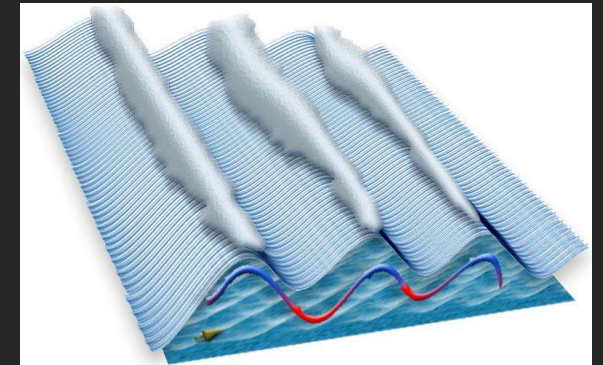
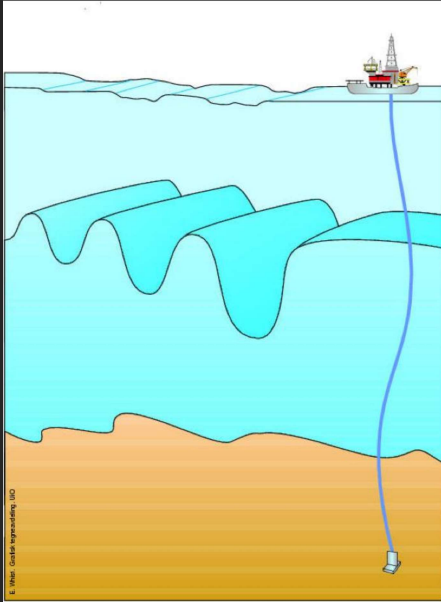
- “Wave of translation” - travel large distances without change of form
- Speed depends on amplitude - larger/faster
- Do not merge - solitary behaviour

# Internal Solitary Waves





# Internal Solitary Waves



- Travel on density interfaces within stably stratified fluids; **balance between nonlinear wave steepening and linear wave dispersion**
- They occur in many geophysical settings including coastal zones & river outflows, estuaries, lakes, reservoirs & fjords, oceans & marginal seas and the atmosphere.
- Gaseous plasmas, liquid crystals, acoustics, solid state physics, optical fibres, Bose-Einstein condensates, etc.

Helfrich & Melville (2006), Lamb (2014), Boegman & Stastna (2019), *Annual Review of Fluid Mechanics*.

# Properties of ISWs

- Wave speeds 0.1-1 m/s; front 10-100 km; length 0.1-1 km, amplitudes 10-100m.
- Rank-ordered wave packet
- $a=120$  m,  $H=340$  m: [Duda et al. 2004](#) (S China Sea)
- $a=240$  m,  $c=2.55$  m/s: [Huang et al. 2016](#) (S China Sea)



Forth Road Bridge (155m)

# Physical Oceanography

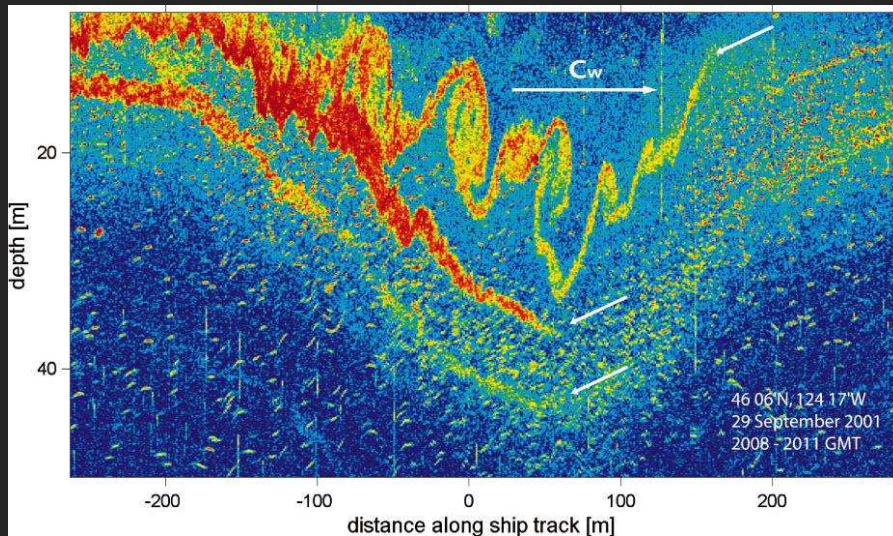


Fig 1. Moum *et al.* 2003. Oregon continental-shelf.

- Source of momentum & mixing
- Vertical transport of heat & nutrients
- Resuspension of sedimentary material
- Transport of mass over large distances.

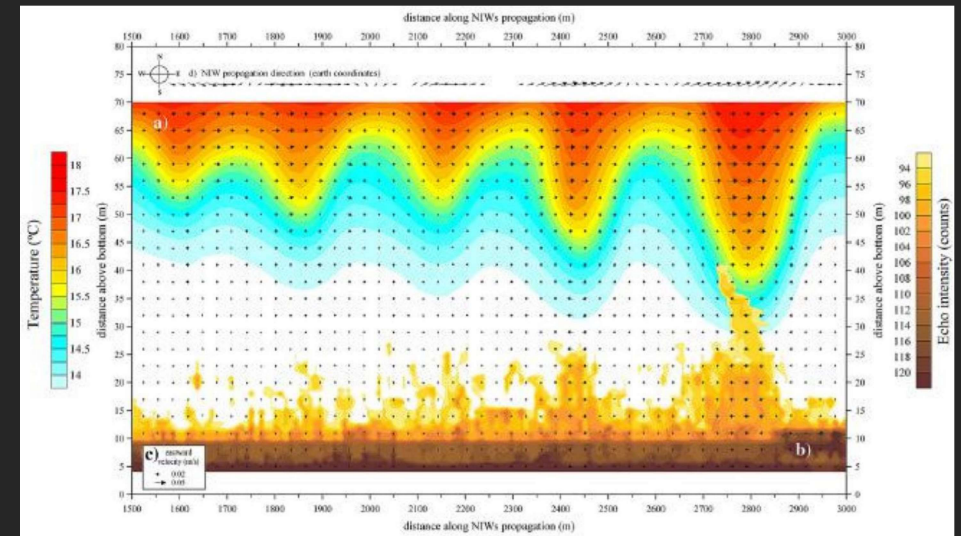
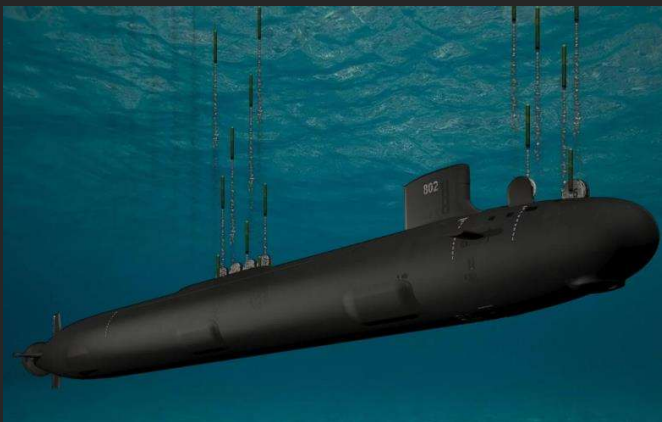
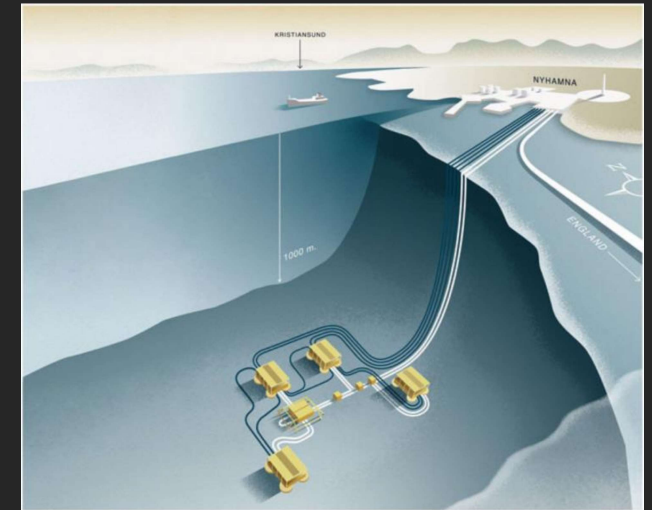
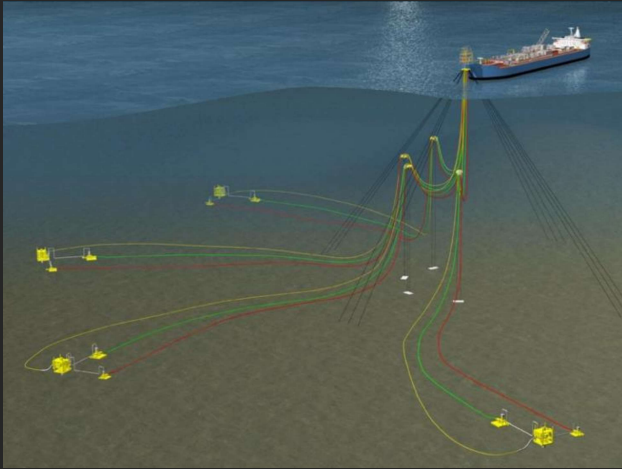


Fig 2. Quesada *et al.* 2007. Western Portuguese mid-shelf.



# Applications





# Remote Sensing of ISWs

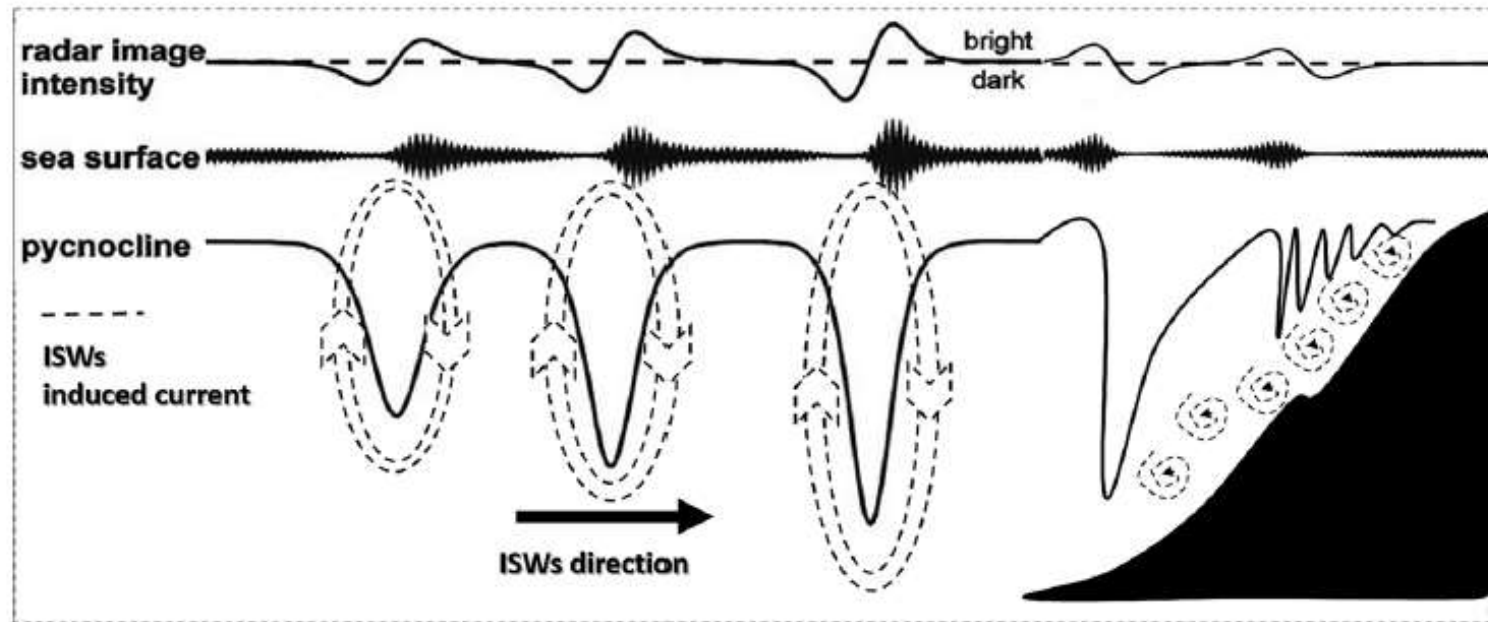
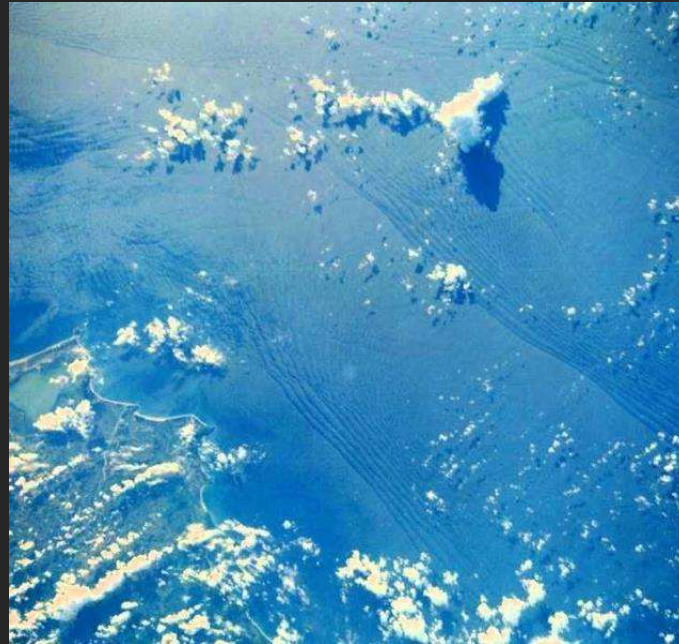
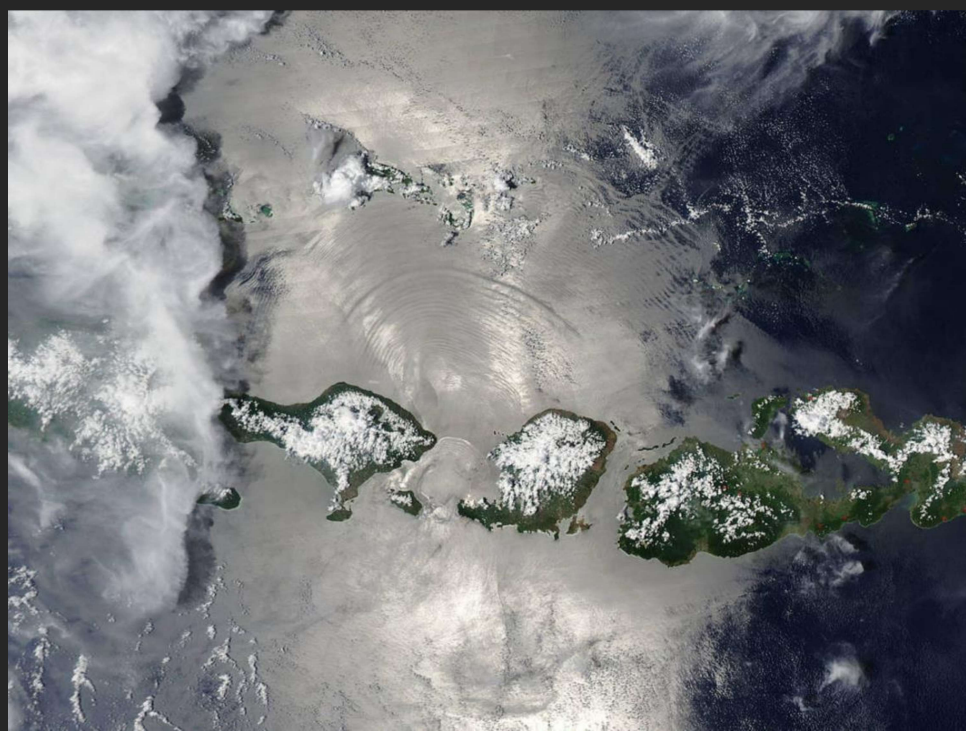


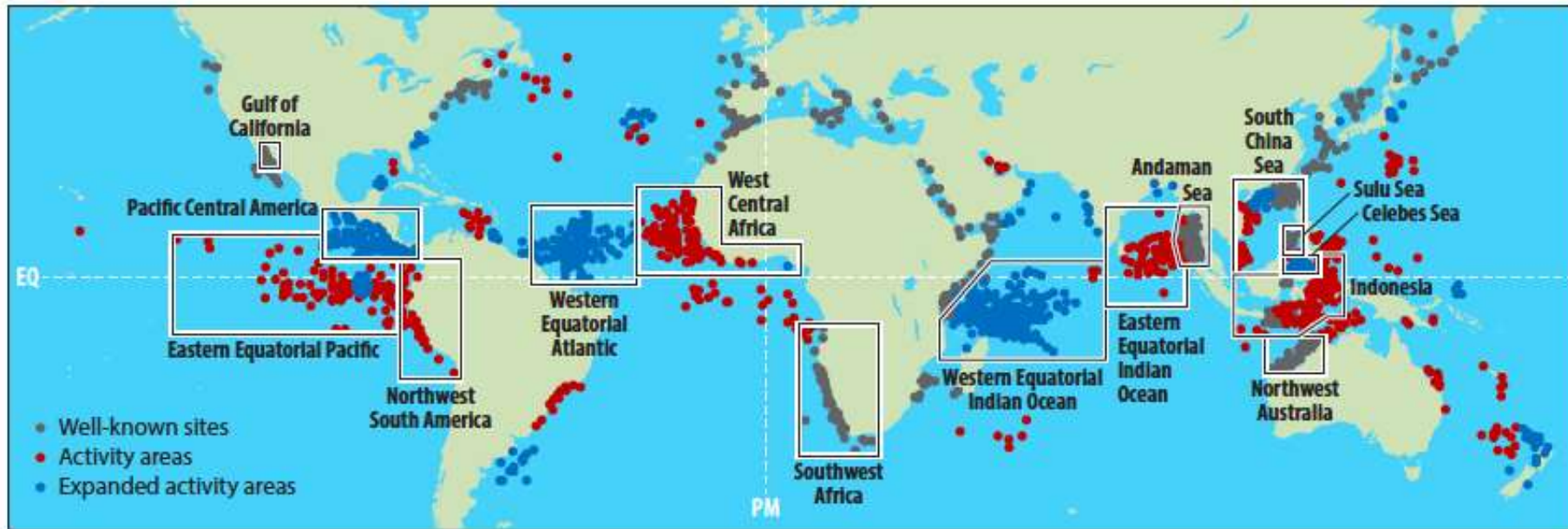
Figure 2. Schematic plot of remote sensing mechanism of internal solitary waves [adapted from earth.esa.int].

# Remote Sensing of ISWs



NASA satellite photos of (i) the Lombok strait, (ii) Hainan Island and (iii) the strait of Gibraltar.

# Global Distribution of ISWs



**Figure 1**

Global distribution of internal solitary waves. Internal waves observed from August 2002 through May 2004. Shown are well-known occurrence sites (*gray*), new areas of activity (*red*), and areas of geographically expanded activity (*blue*). Figure adapted from Jackson (2007), with permission from Wiley.

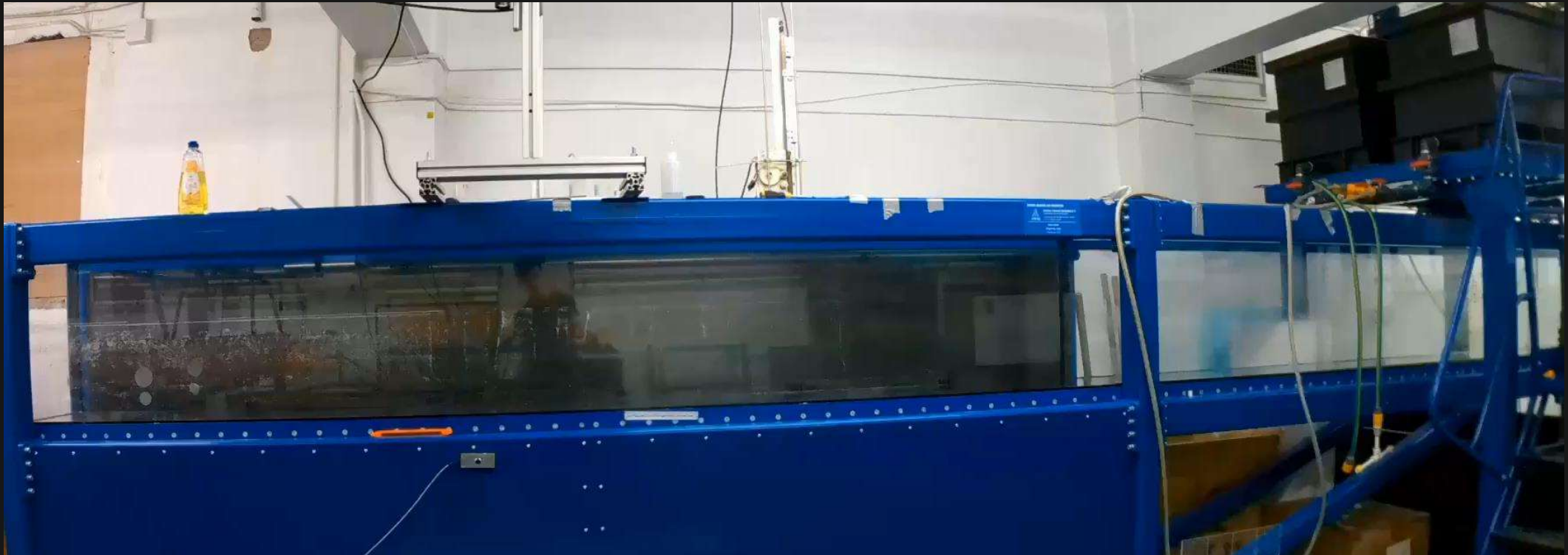
- Jackson (2007) observed ISW packets on nearly every coastline worldwide on 95% of observation days. Image taken from Boegman & Stastna (2019), *Annual Review of Fluid Mechanics*.

How do we model Internal Solitary Waves ?



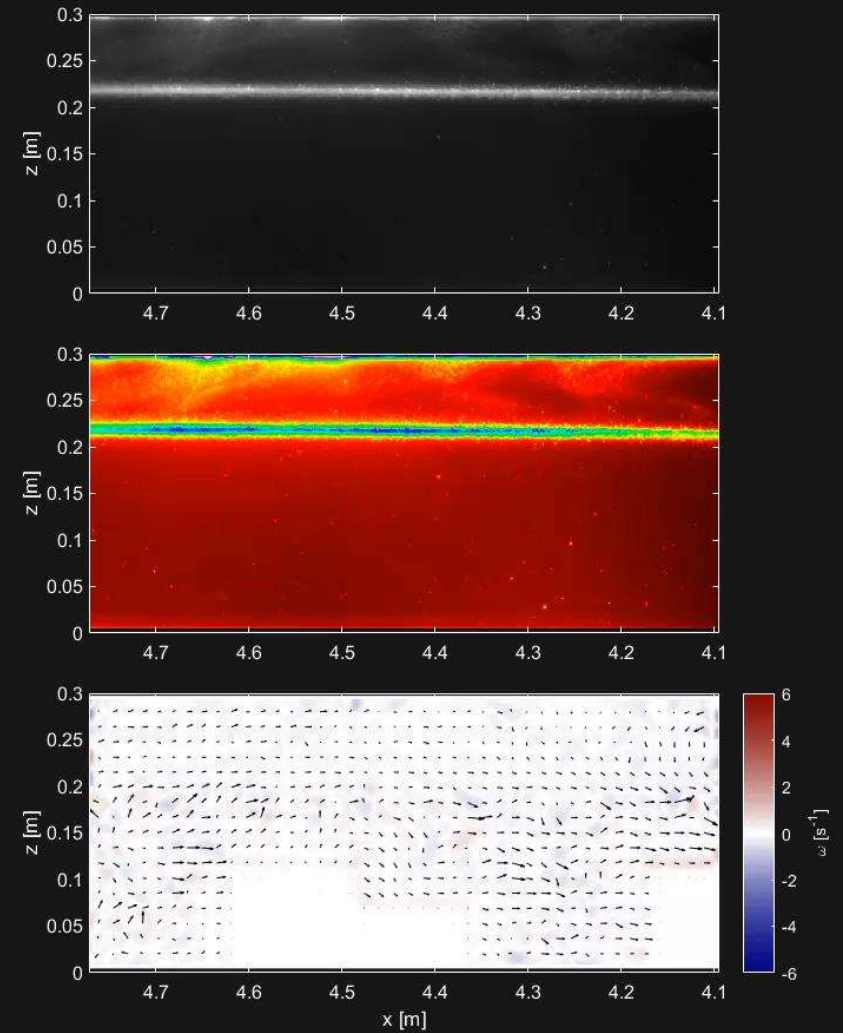


# Experimental Methods



# Experimental Methods

- Particle Image Velocimetry (PIV) – synoptic velocity field.
- Microconductivity Sensors – traverse of the density field.



# Numerical Model – Spectral Parallel Incompressible Navier-Stokes Solver (SPINS)

$$\frac{\partial \vec{u}}{\partial t} + \vec{u} \cdot \vec{\nabla} \vec{u} = -\frac{1}{\rho_0} \vec{\nabla} P + \nu \nabla^2 \vec{u} - \frac{\rho g}{\rho_0} \hat{k}$$

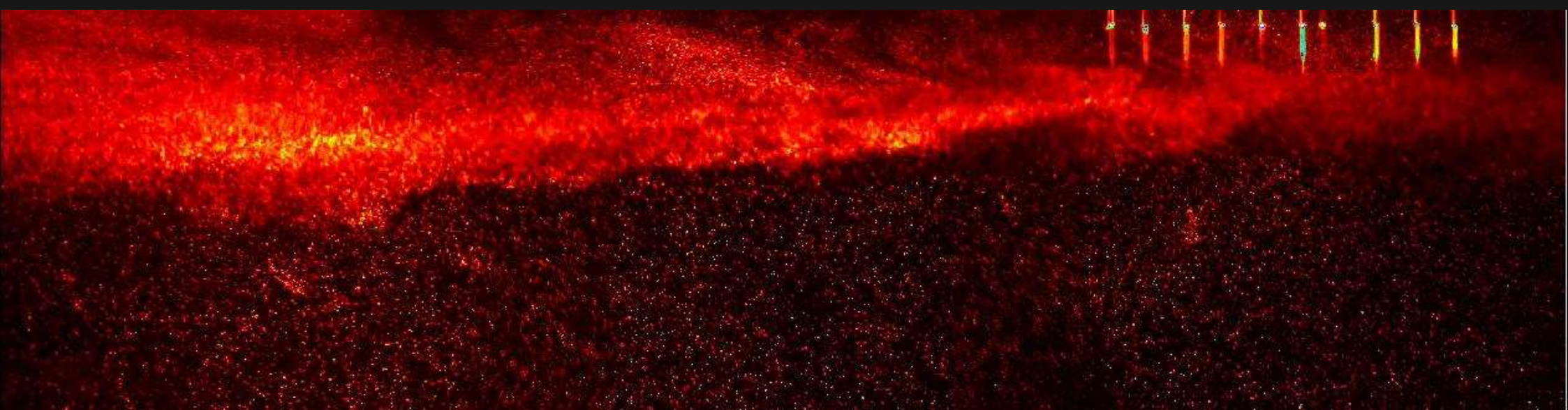
$$\nabla \cdot \vec{u} = 0$$

$$\frac{\partial \rho}{\partial t} + \vec{u} \cdot \vec{\nabla} \rho = \kappa \nabla^2 \rho$$



# Some Results

# Laboratory Observation of an unstable ISW



- Fructus, Carr, Grue, Jensen & Davies. *J. Fluid Mech.* (2009).
- Carr, Franklin, King, Davies, Grue & Dritschel. *J. Fluid Mech.* (2017).

# Numerical Results

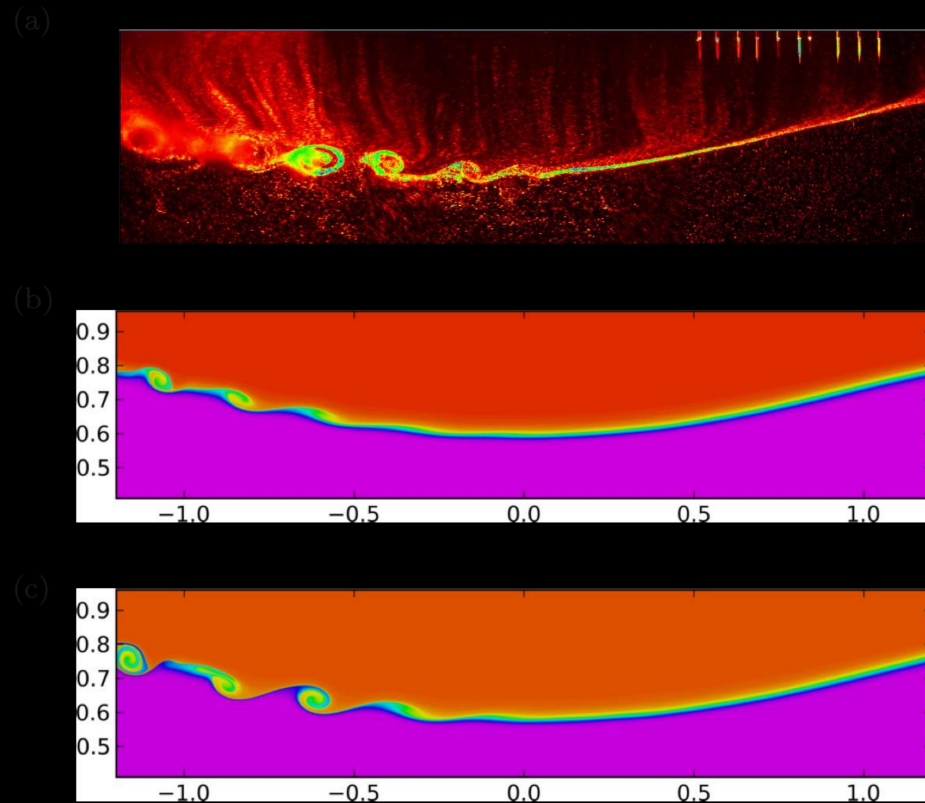
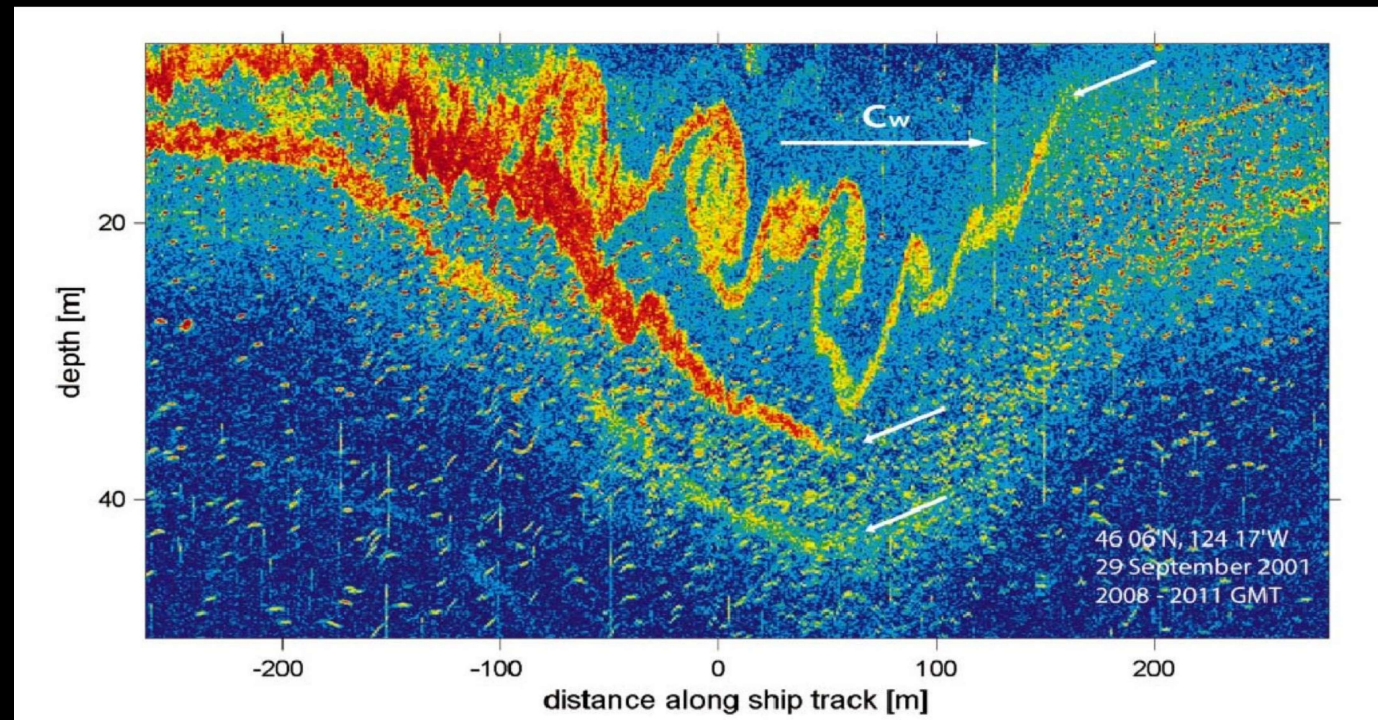


Fig: (a) Experimental image and (b) corresponding numerically computed buoyancy field when  $a_{\text{lower}}$  and (c)  $a_{\text{upper}}$  are matched with the laboratory wave depicted in (a).

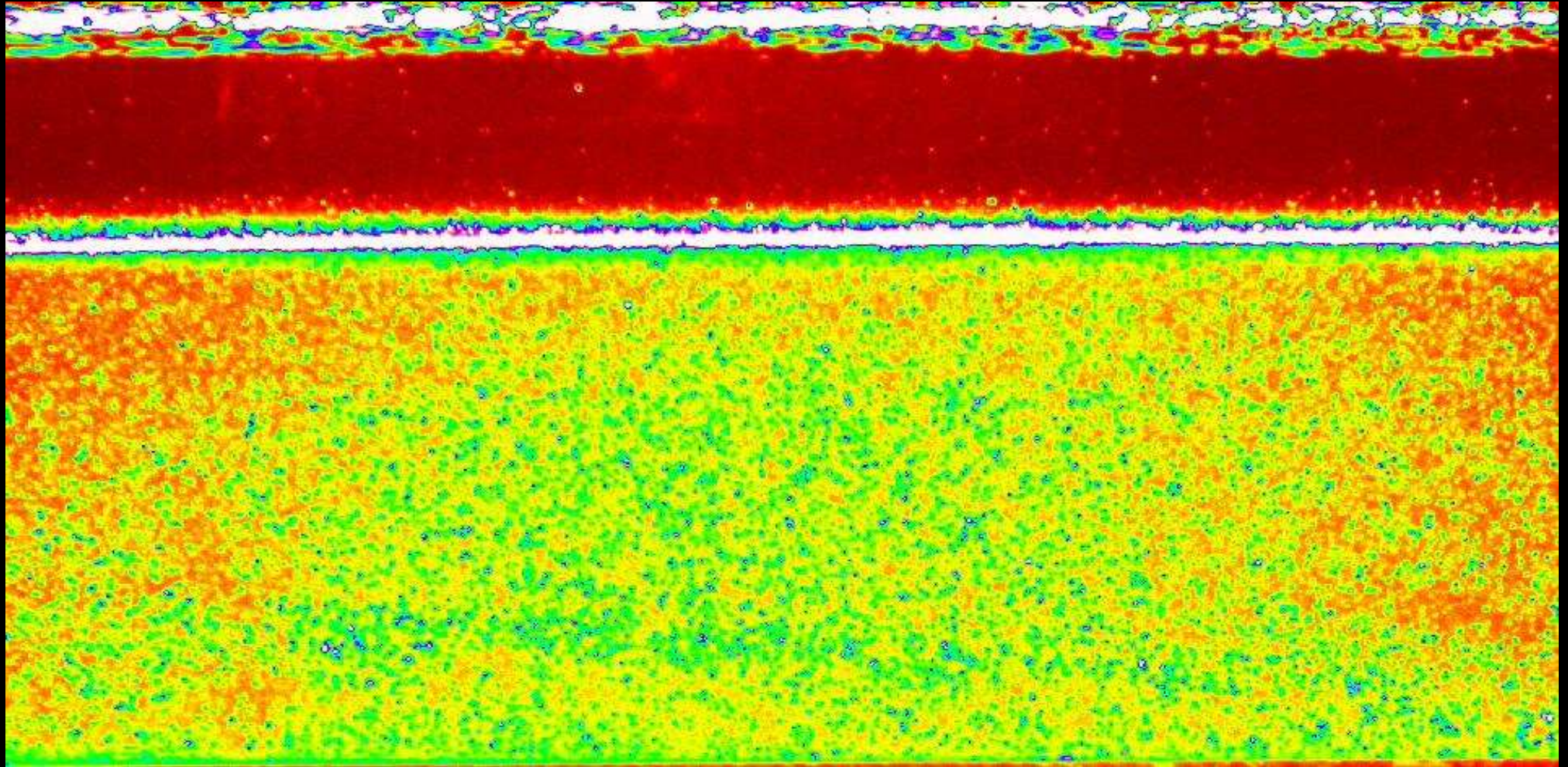
# Field Observation



Moum *et al.* 2003. Oregon Continental Shelf.

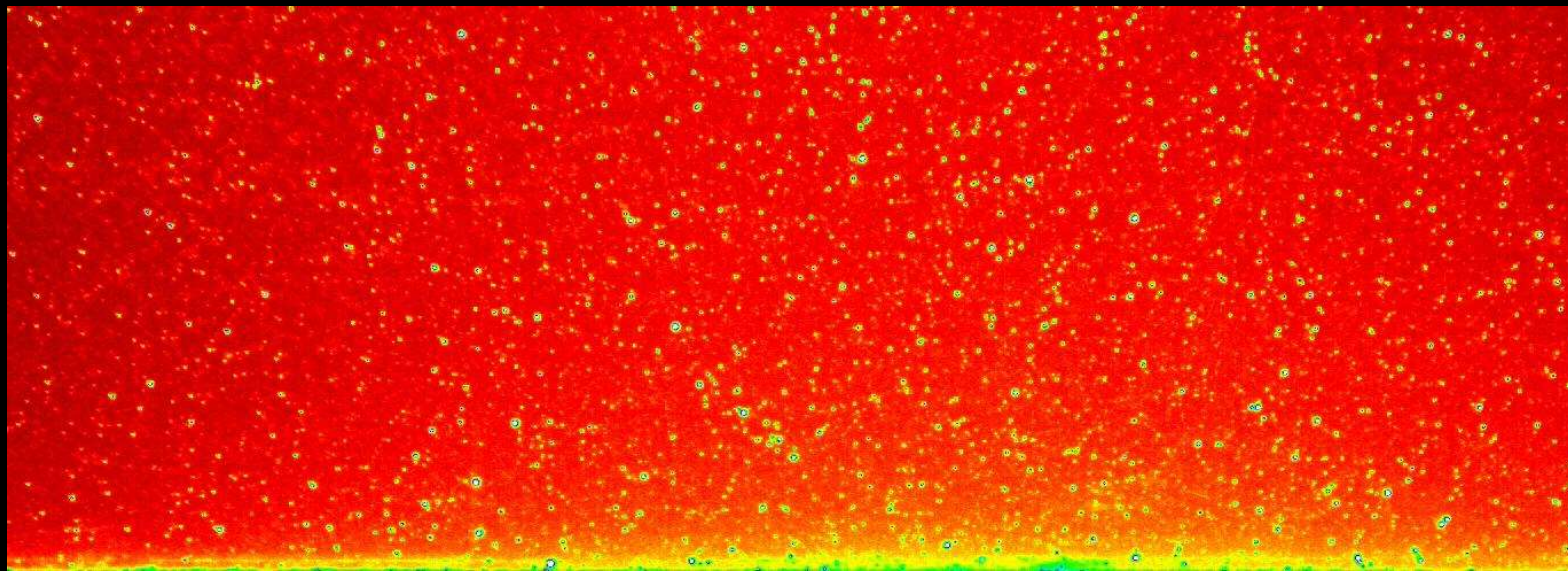
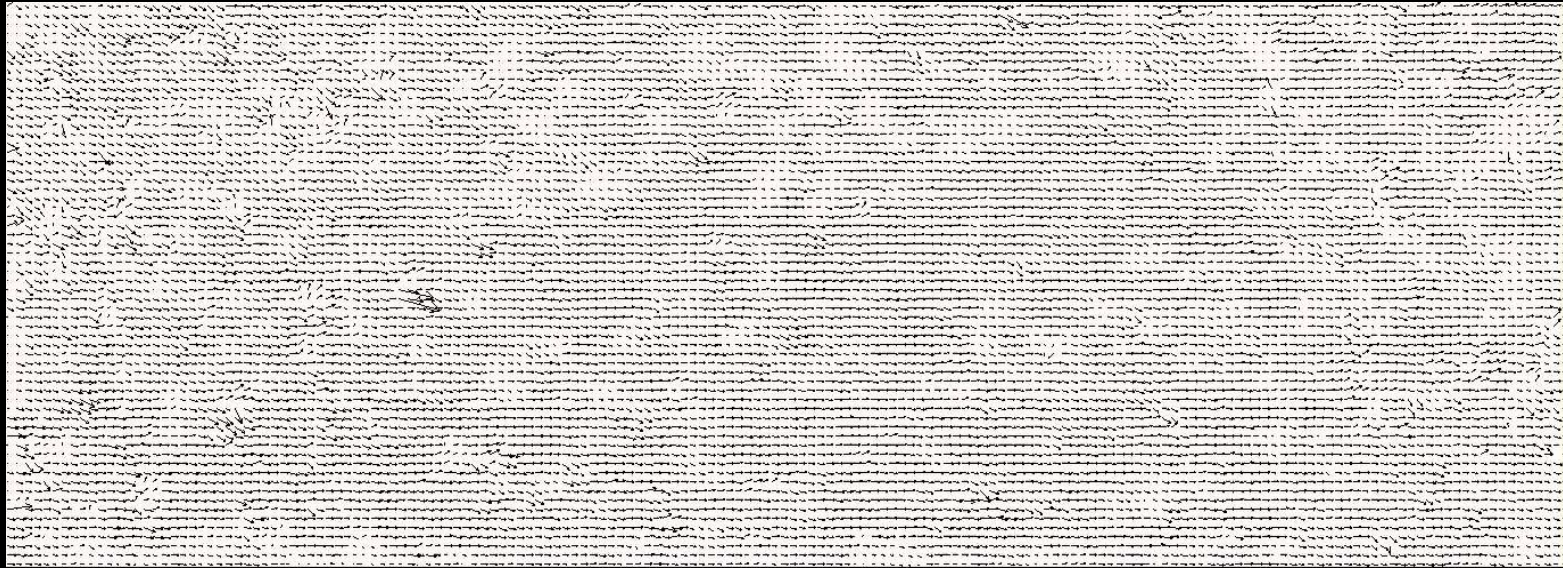


**Carr & Davies. *Phys. Fluids*. (2006)**



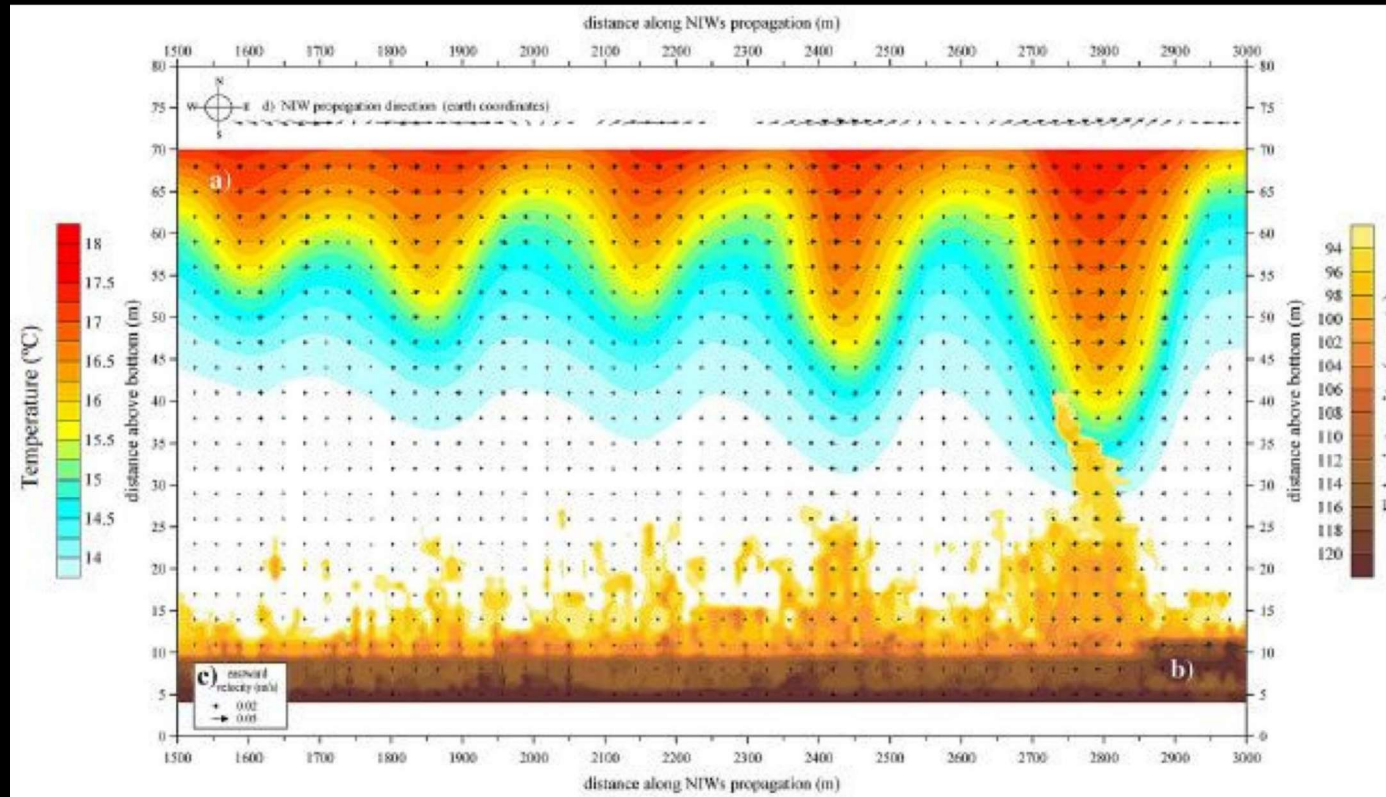


**Carr, Davies & Shivaram. *Phys. Fluids*. (2008)**



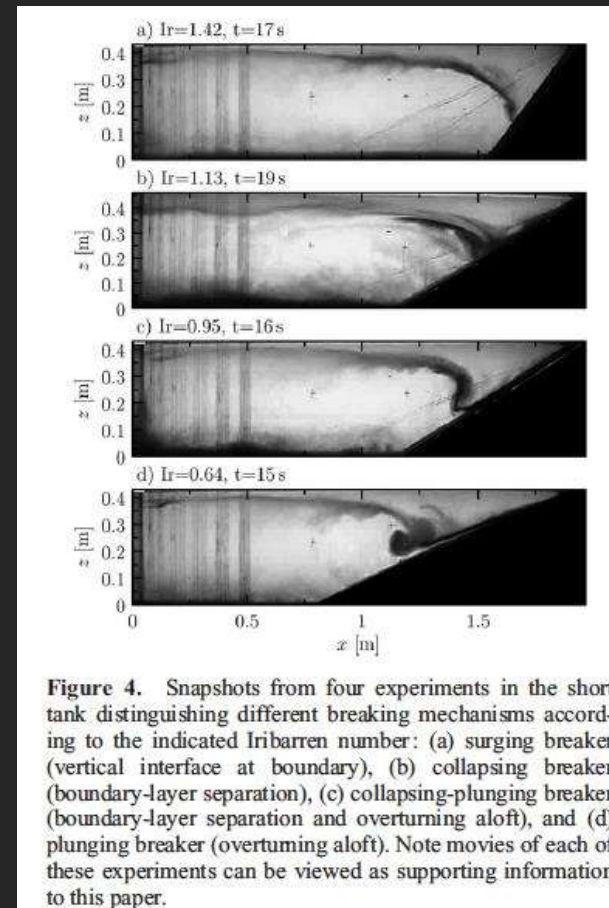
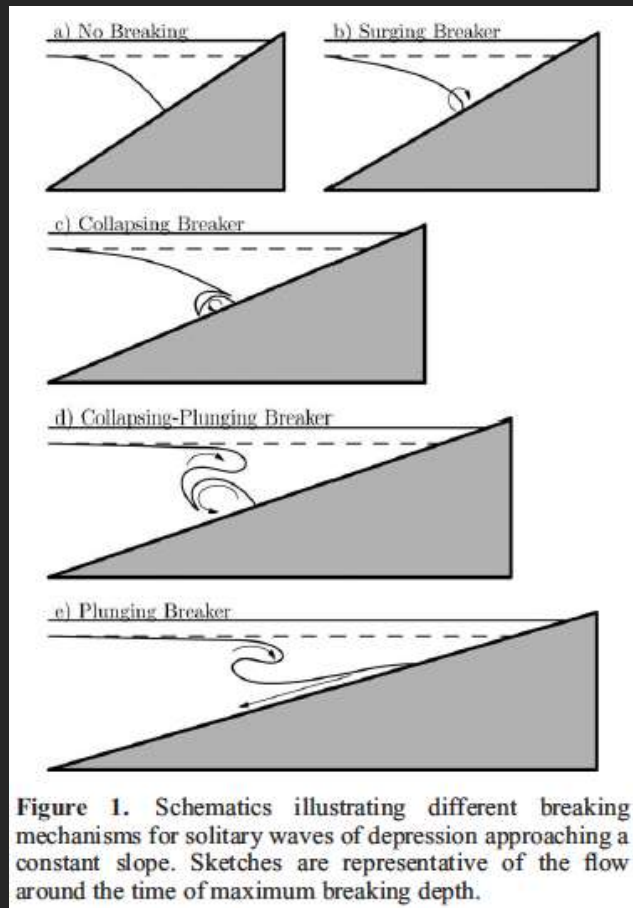


# Field Observation



Quaresma *et al.* 2007. Western Portuguese Mid-shelf.

# Classification of shoaling breakers



- Sutherland *et al.* 2013. JGR.

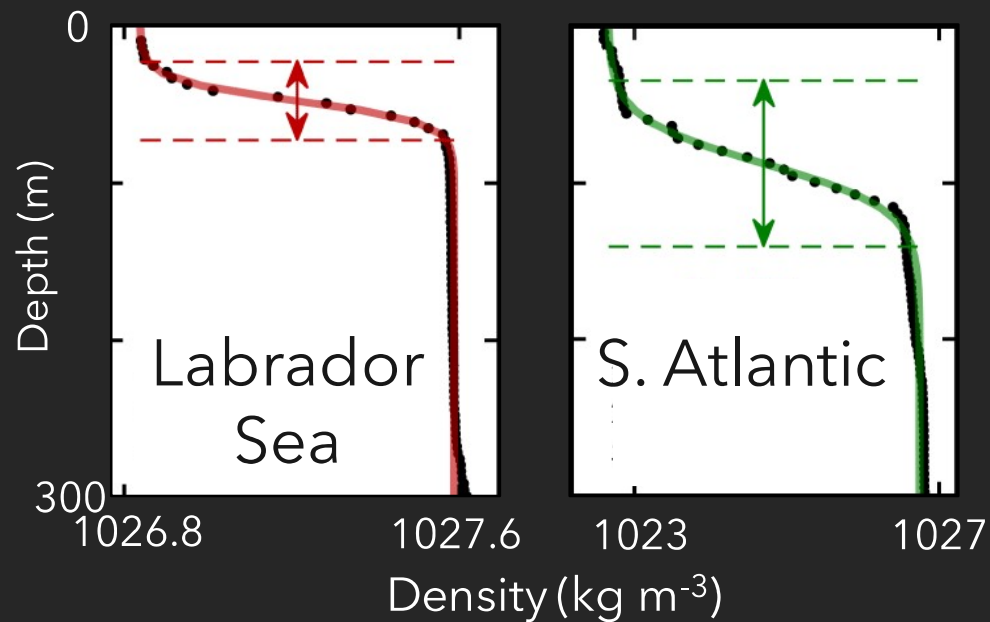


# Classification of shoaling breakers

$$Ir = \frac{s}{\sqrt{a/\lambda}} \quad \text{Iribarren Number}$$

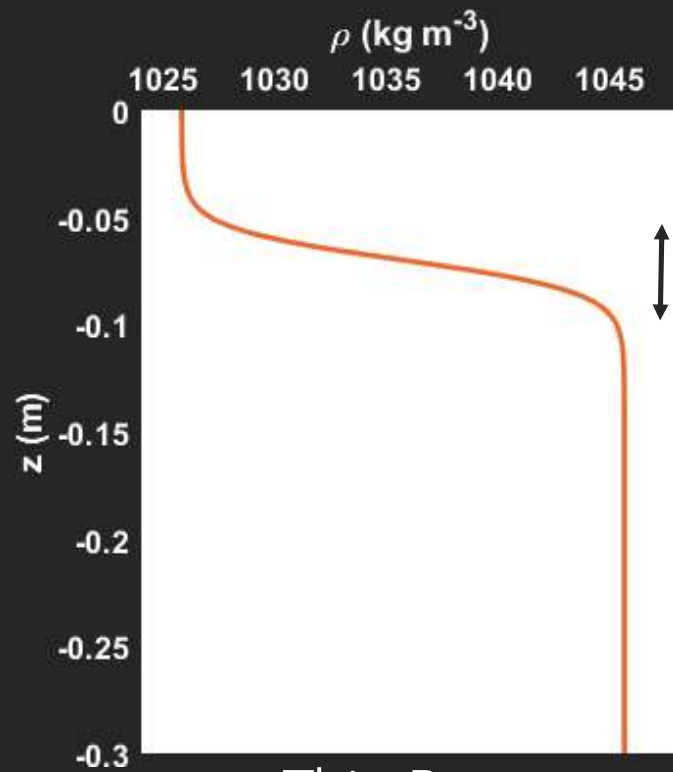
- Qualitative change from plunging, to collapsing to surging as  $Ir$  increases (Boegman *et al.* 2005)
- On gentle slopes fission followed by bolus formation (Aghsaei *et al.* 2010)

# Stratification in the Ocean

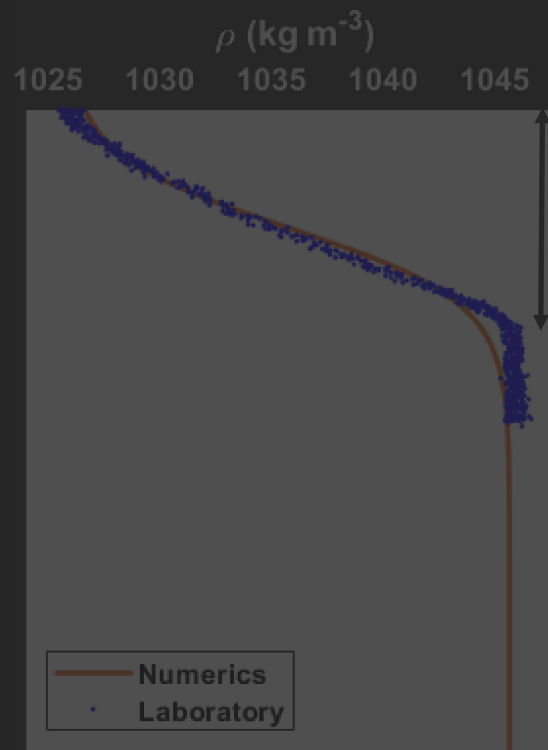


$$\rho(z) = \rho_2 + \Delta\rho \tanh\left(\frac{z - z_{pyc}}{h_{pyc}}\right)$$

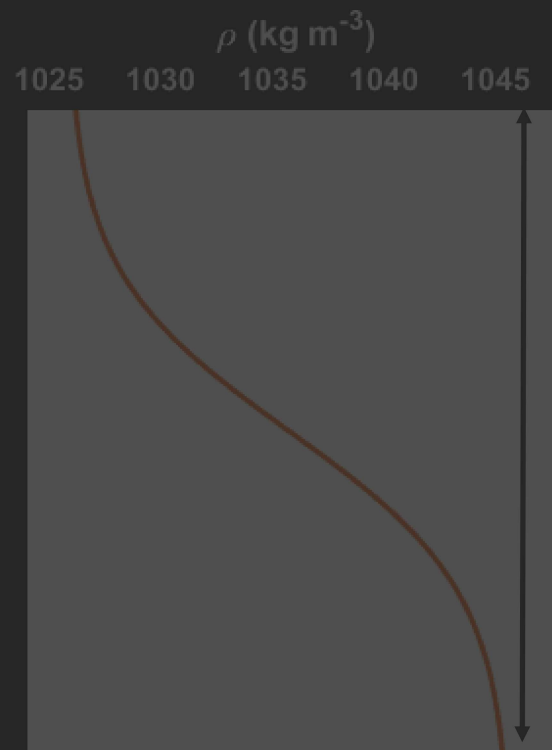
# Stratification Investigated



Thin Pyc.

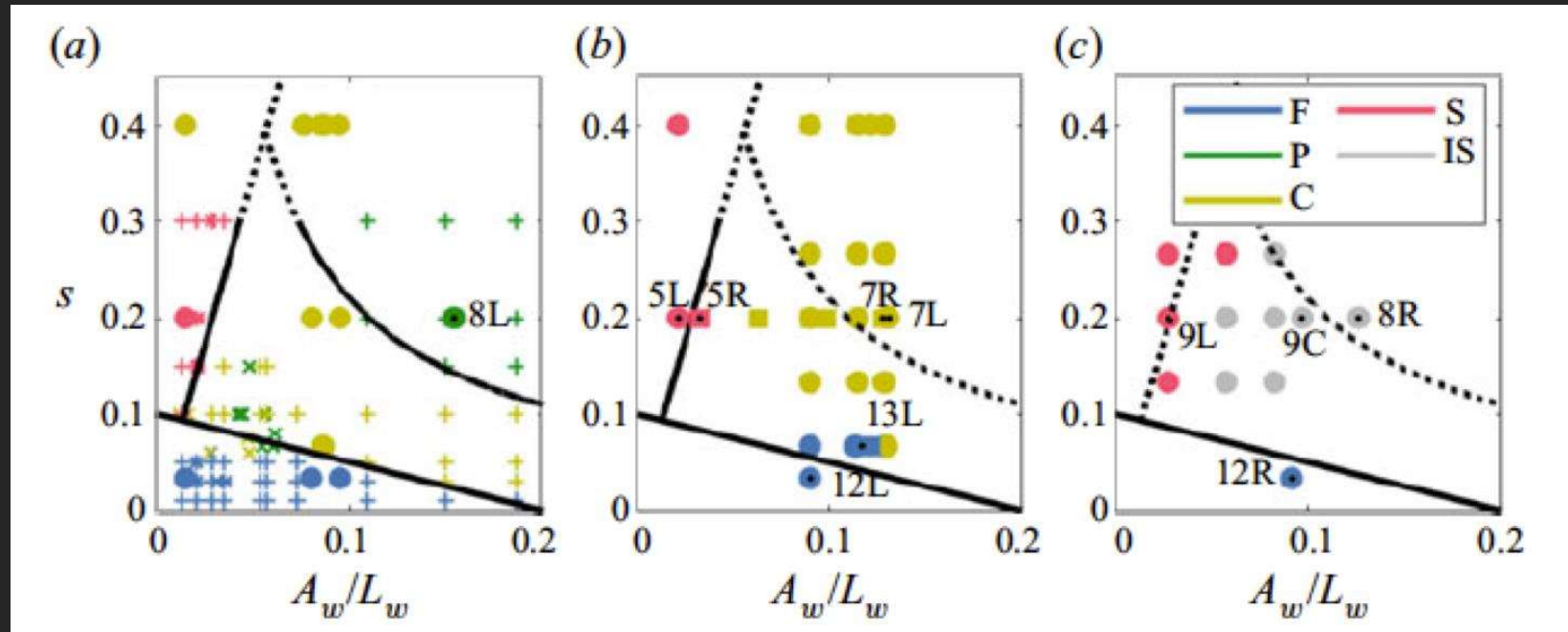


Surface Stratification



Broad Stratification

# Shoaling Classification



(a) thin tanh profile, (b) surface stratification and (c) broad tanh profile

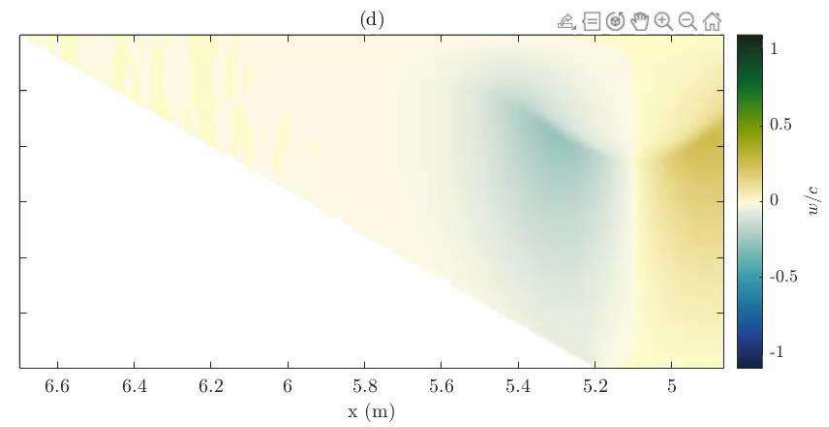
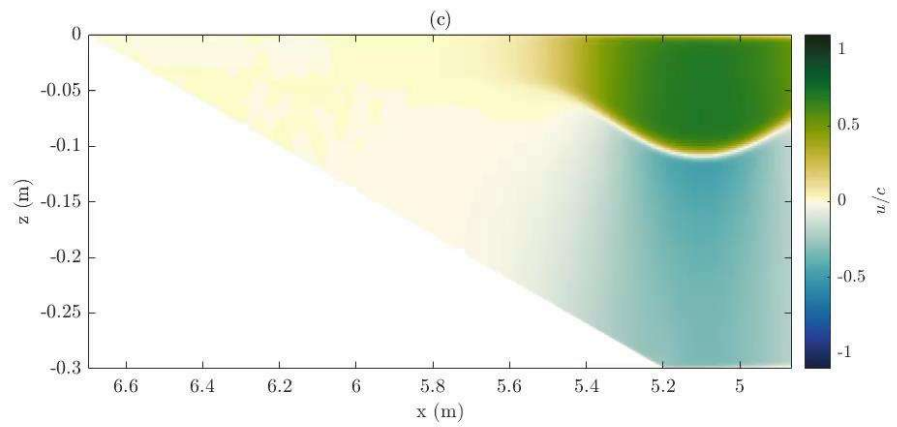
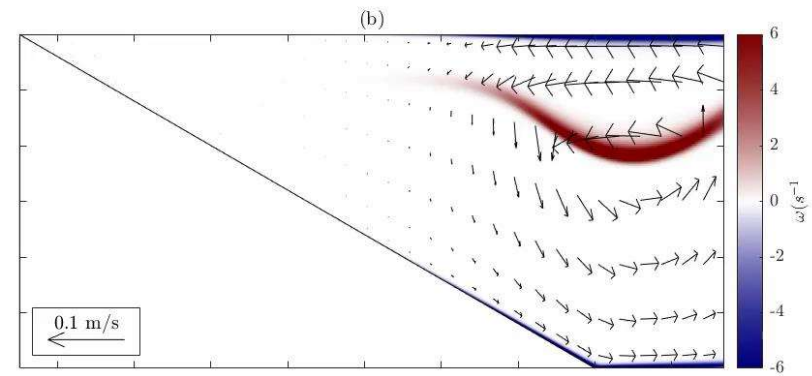
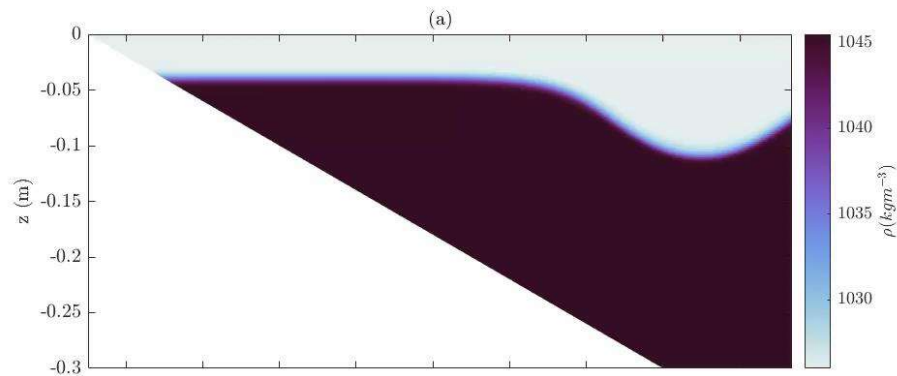
(+) Aghsaei et al., 2010; *J. Fluid Mech.*

(x) Nakayama et al., 2019; *Phys. Rev. Fluids.*

Filled circle - our numerical, Filled squares - our laboratory

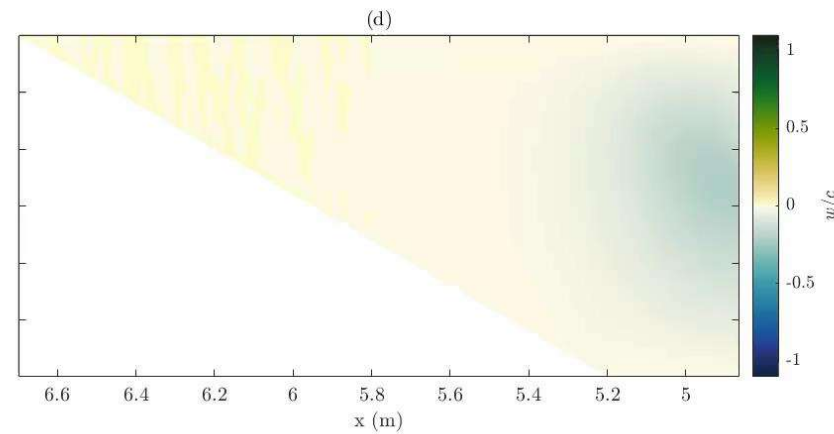
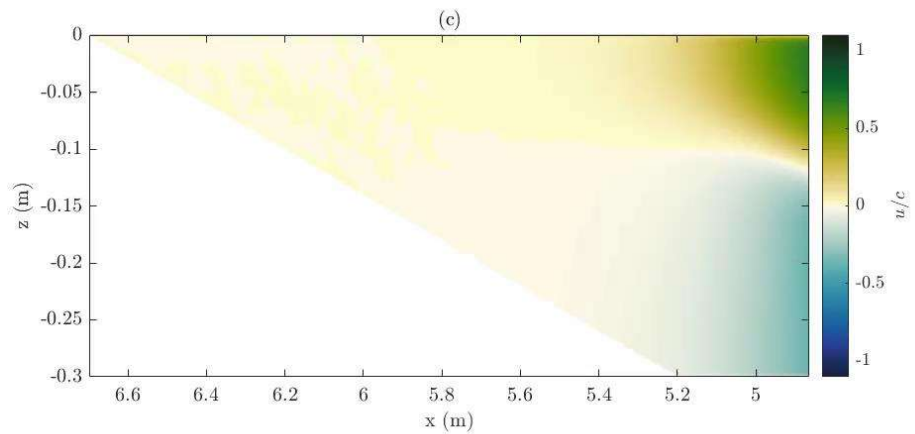
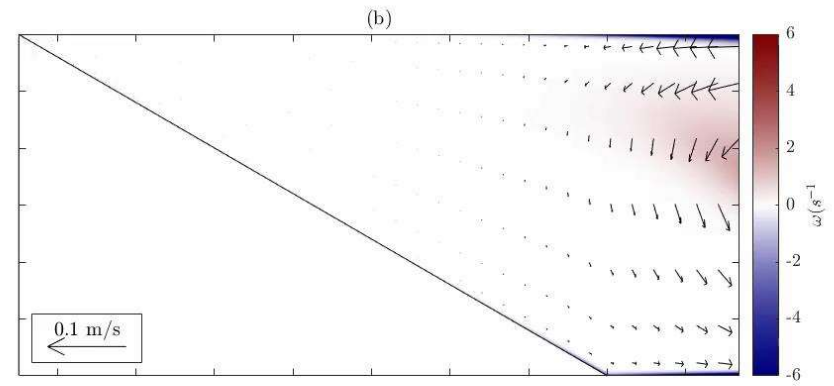
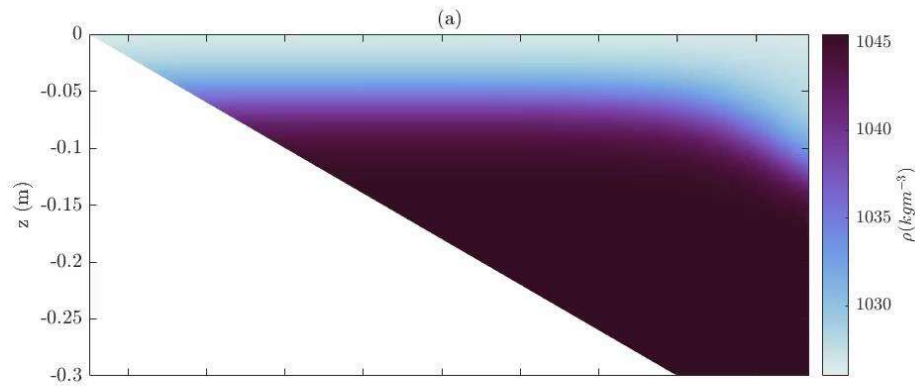


# Plunging example (thin tanh profile - $8L$ )



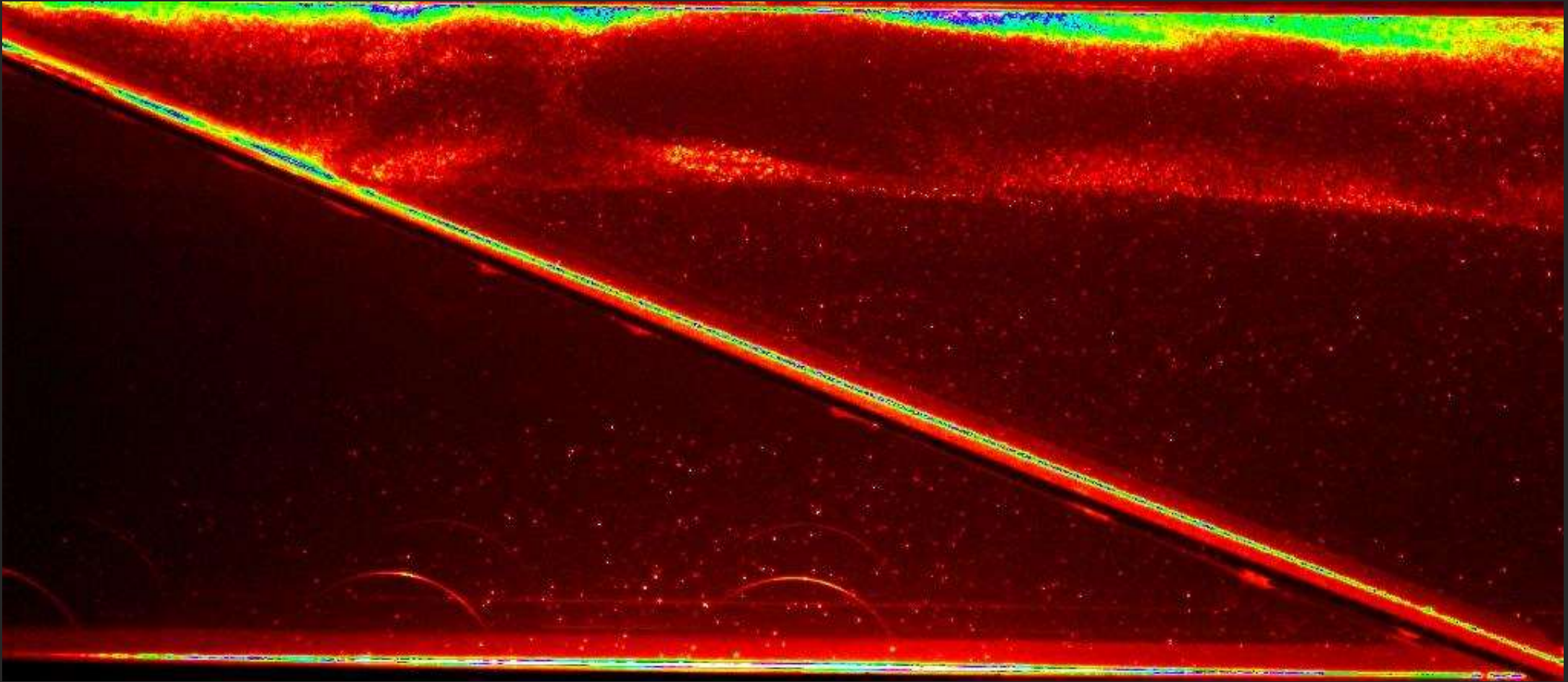
- Overturning leeward face becomes gravitationally unstable

## Collapsing example (surface stratification - 7L)



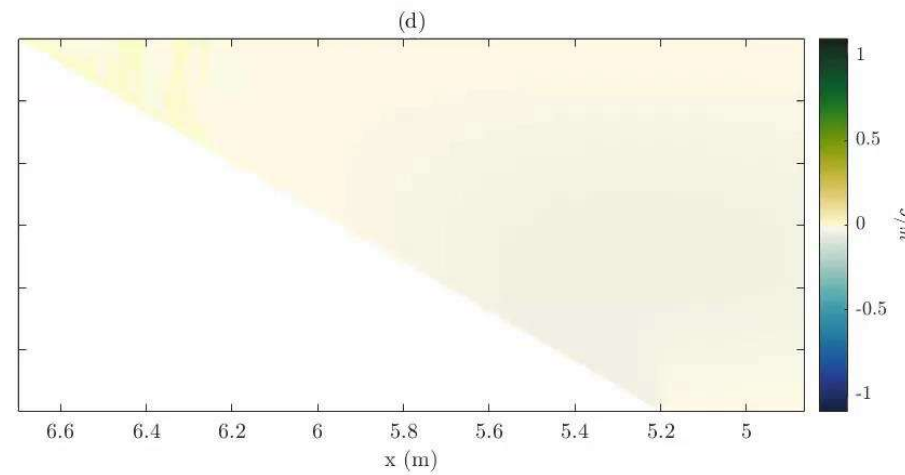
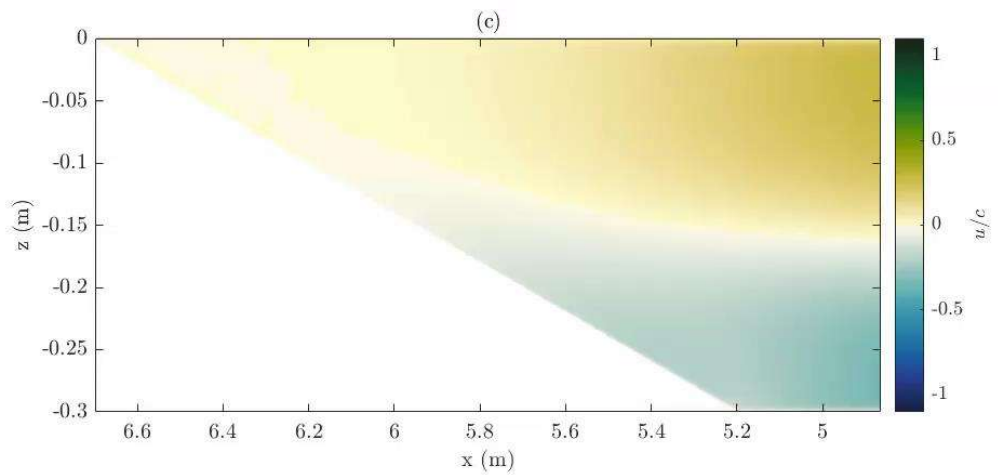
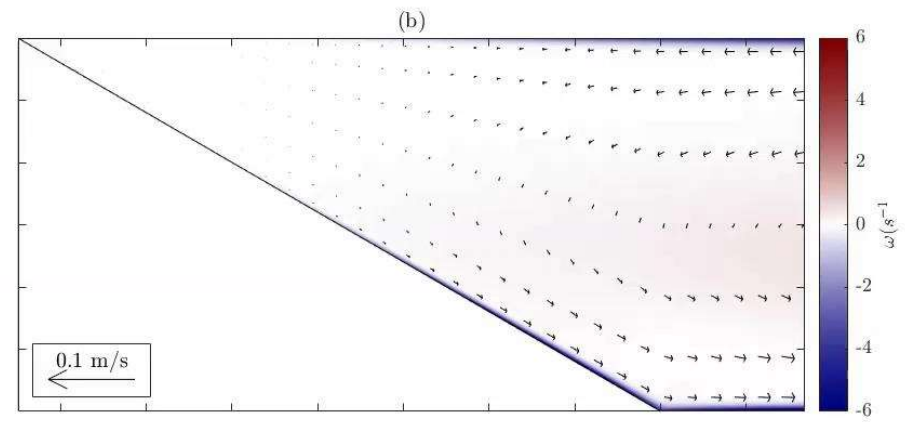
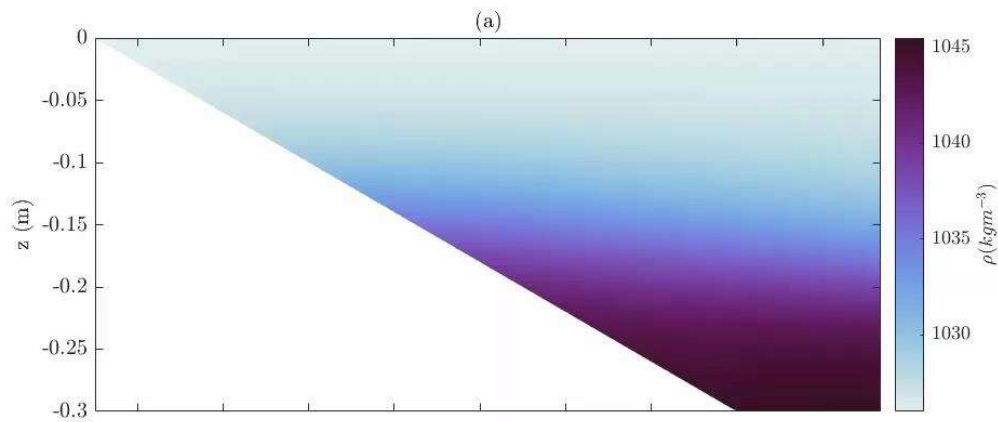
- Separated flow pushes the leeward face backward upon itself
- Density gradient in upper layer suppresses plunging

## Collapsing example - laboratory



Slope height = 0.6m, Generating Volume = 30l

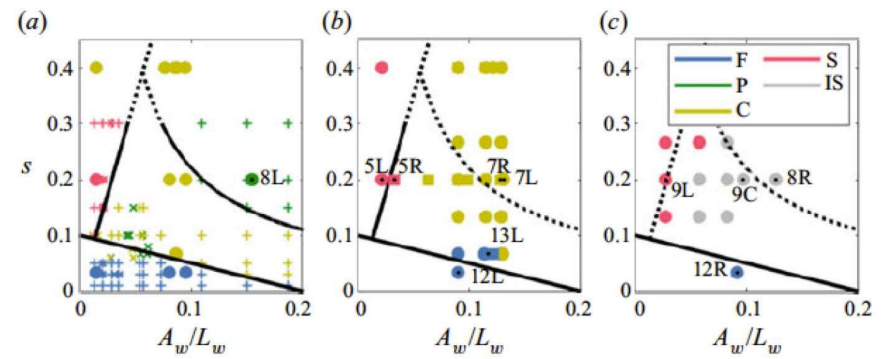
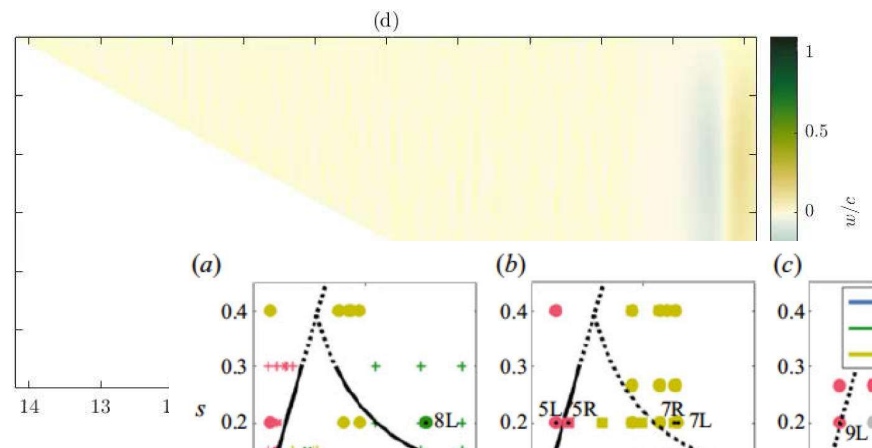
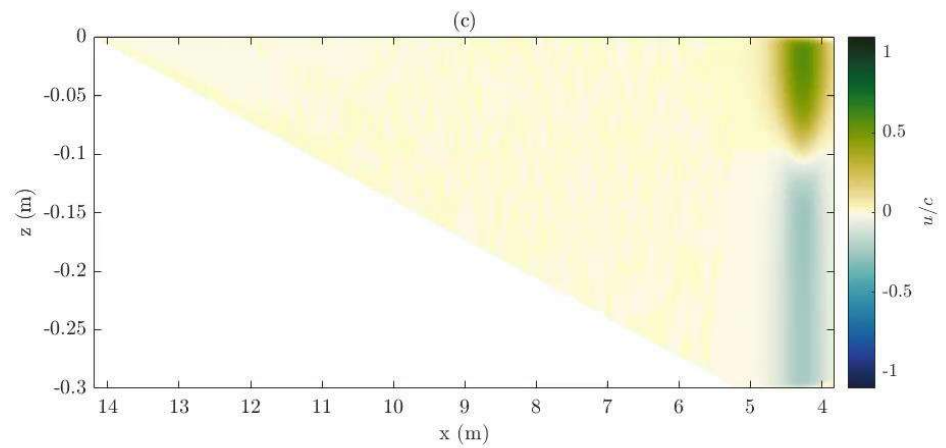
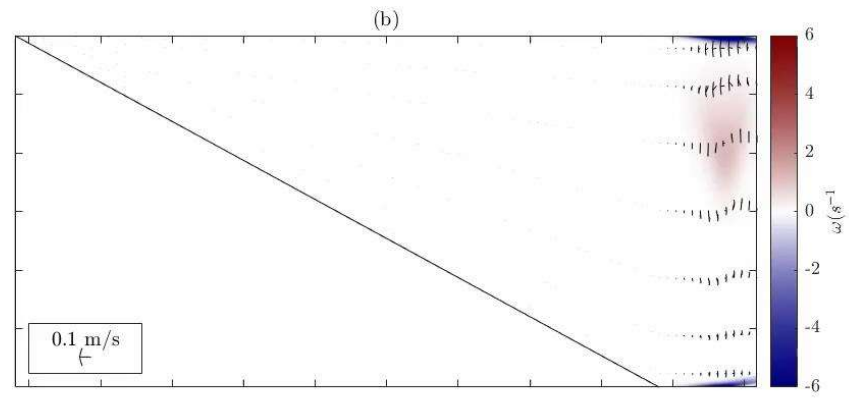
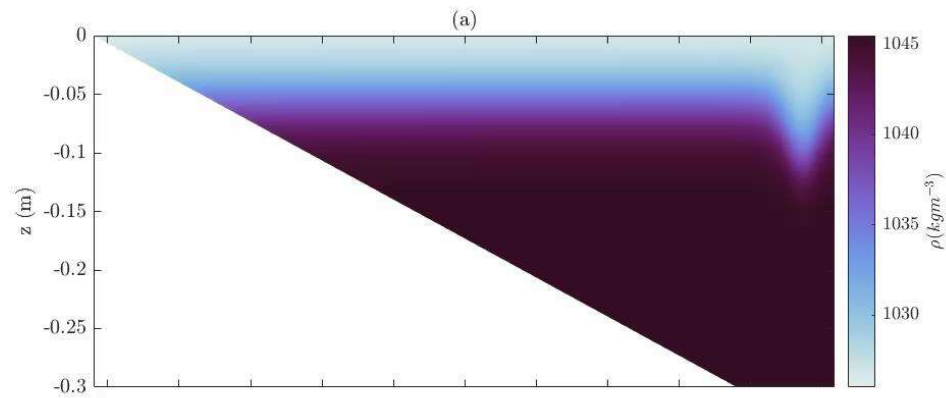
# Surging example from broad tanh profile - 8R



- Density gradient in lower layer suppresses BL vortex formation

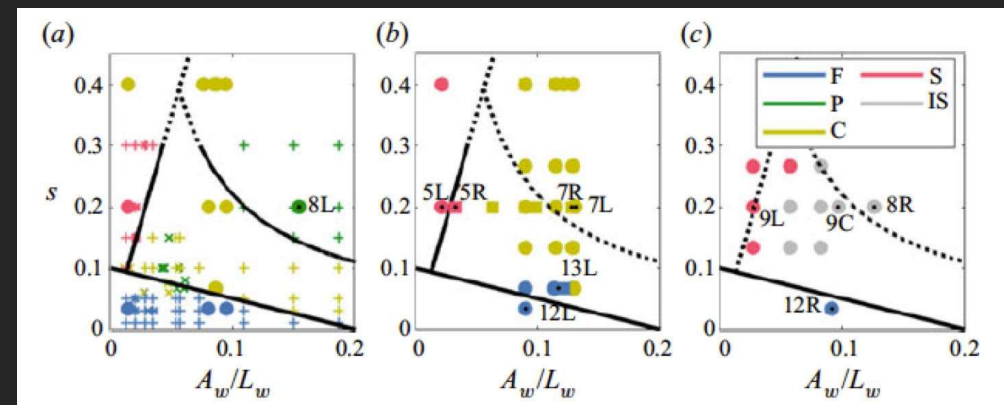


# Fission example (Surface stratification - 12L)

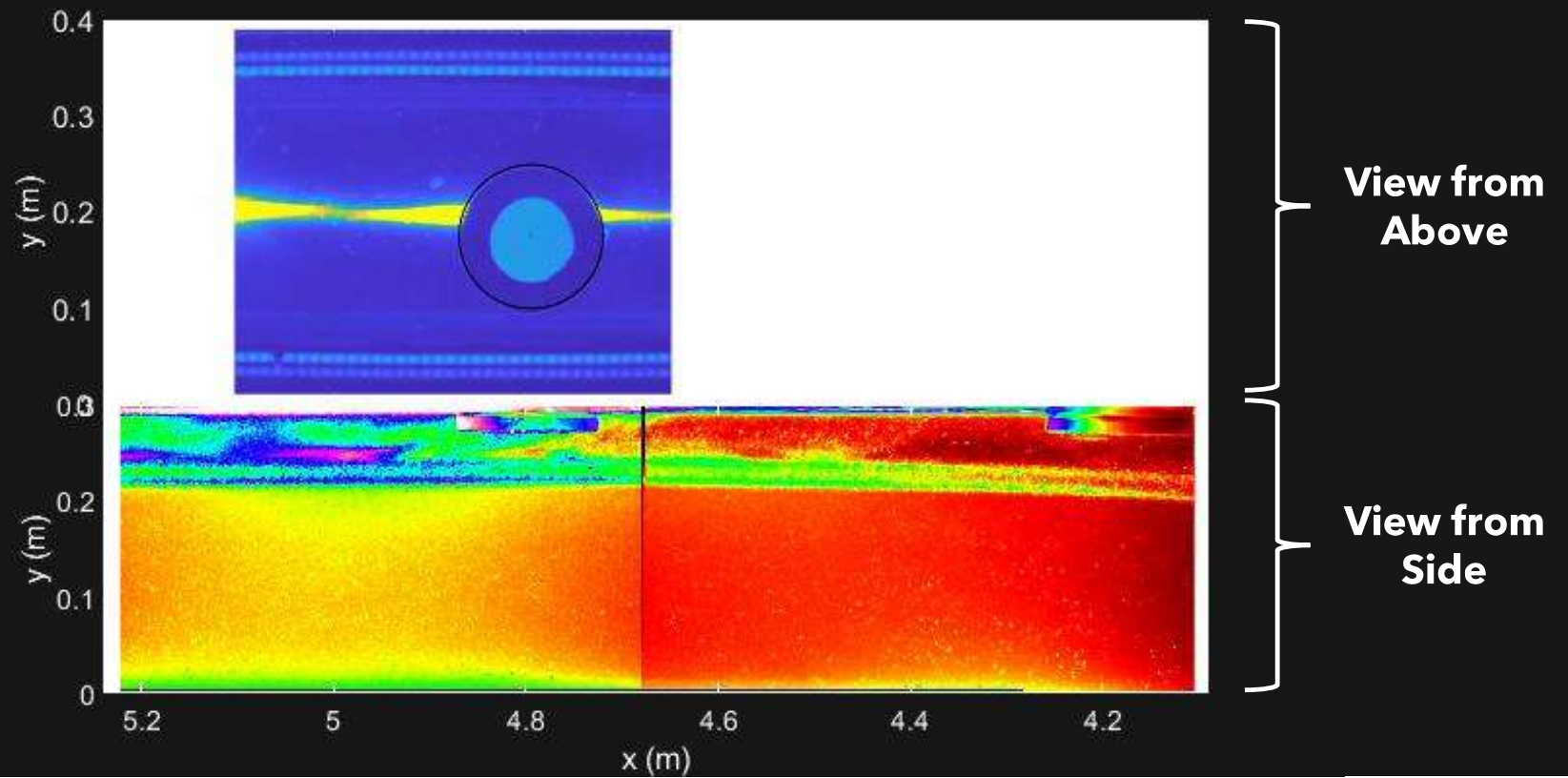


# Main Findings

- Stratification form affects shoaling classification
- Density gradient in the upper layer suppresses plunging dynamics
- Density gradient in the lower layer suppresses collapsing dynamics
- Surging boluses can support K-H instabilities
- On slope steepness representative of continental slopes ( $s=0.03-0.07$ ) we expect fission regardless of the stratification form.



# Ice Motion



## Equation for Ice Motion

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$$F_D = \frac{1}{2} \rho_1 v^2 C_d A$$

$$v = u_x^t - U_{ice}^t$$

$$a = \frac{F_D}{m}$$

$$a_{ice}^t = \frac{1}{2} C_d \frac{\rho_1}{\rho_f} (u_x^t - U_{ice}^t) |u_x^t - U_{ice}^t|$$

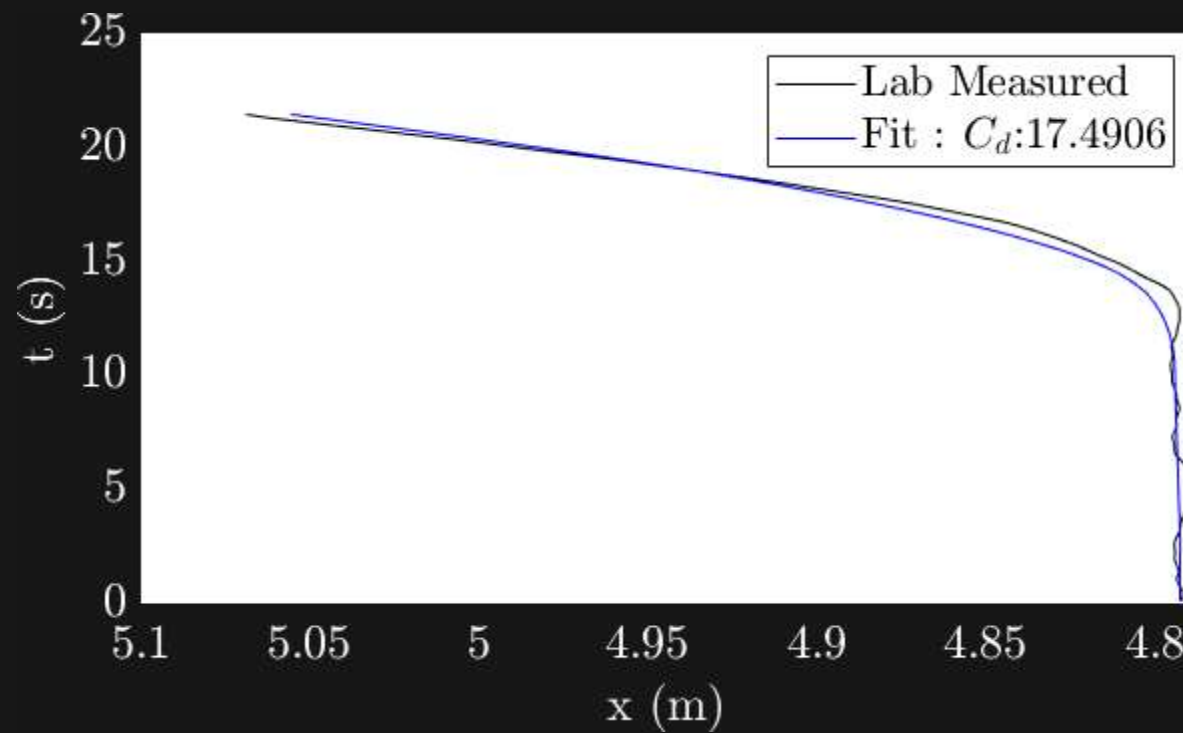
$$U_{ice}^{t+1} = U_{ice}^t + \Delta t a_{ice}^t$$

$$X_{ice}^{t+1} = X_{ice}^t + \Delta t U_{ice}^t$$



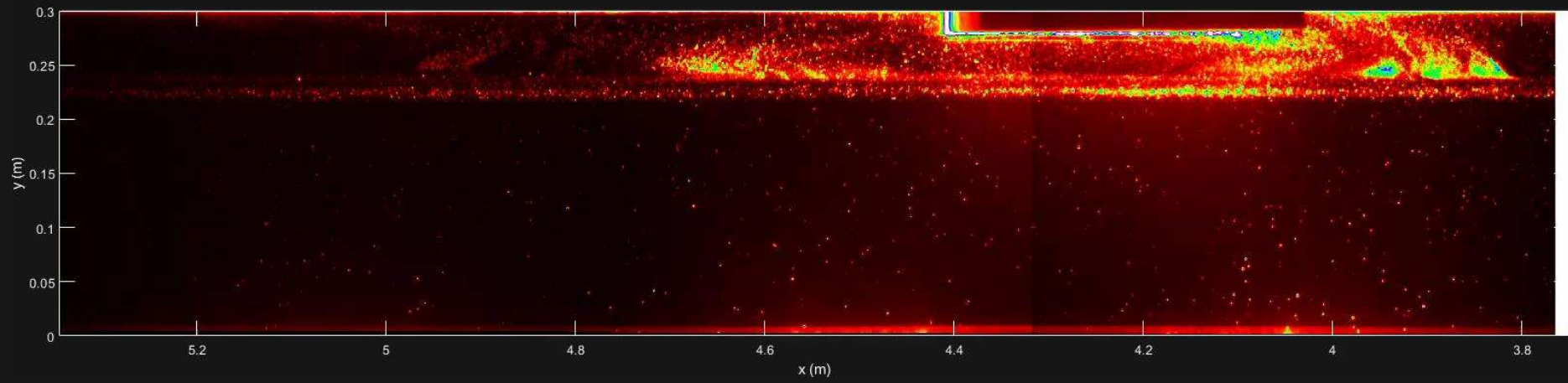
# Calculated Drag Coefficient

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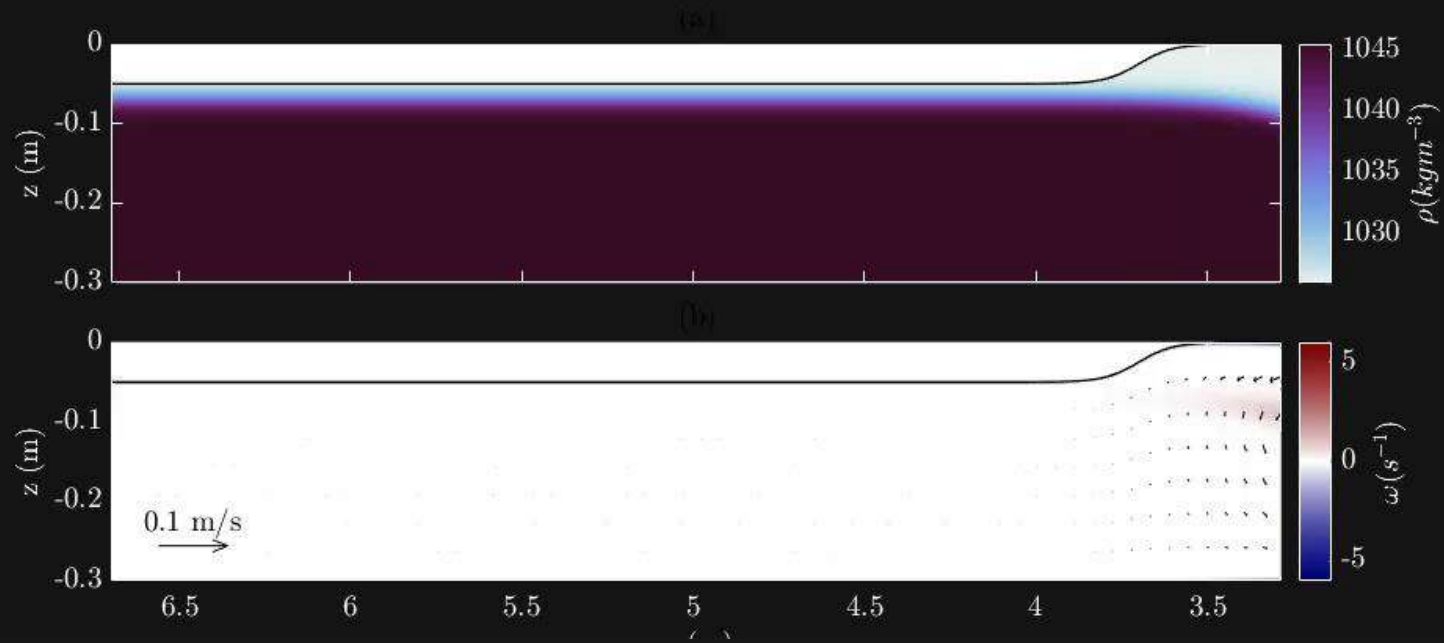


# Large "Ice" Motion

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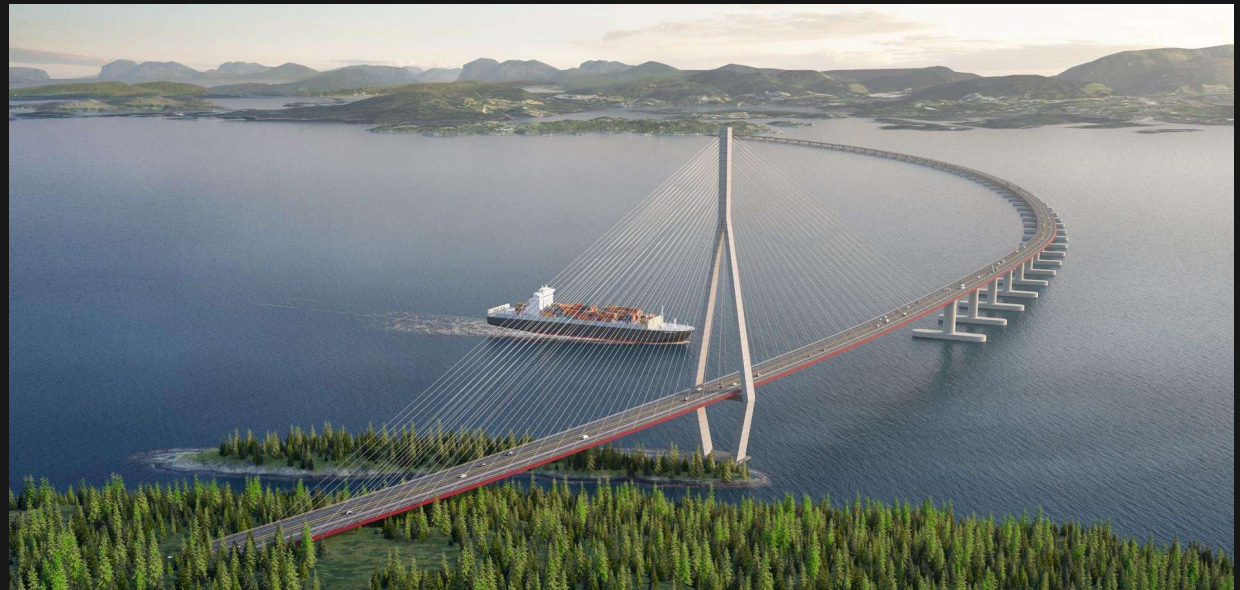
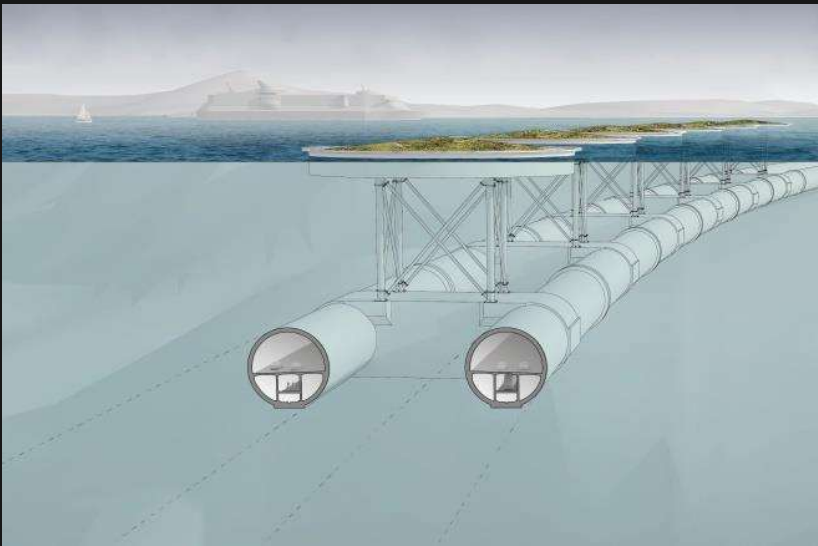
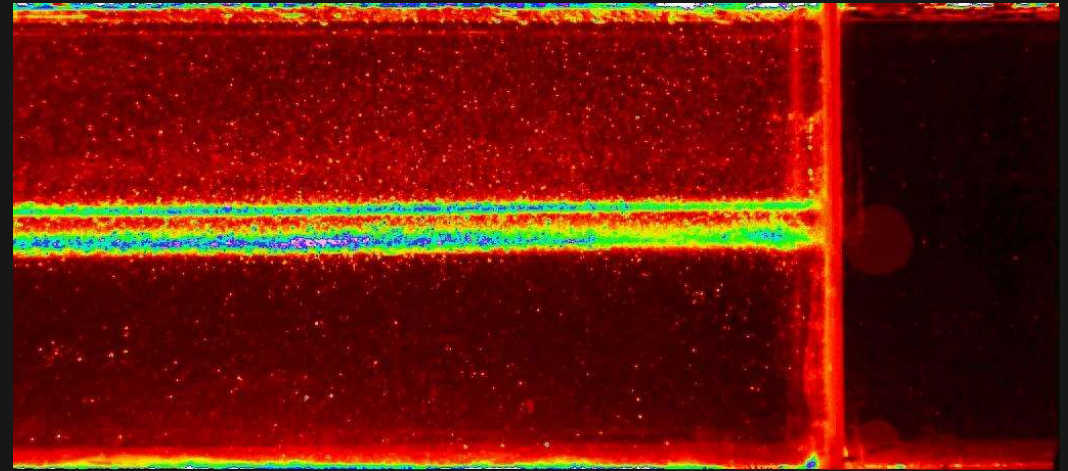


# Numerical Wave/Ice Collisions



# Future interests

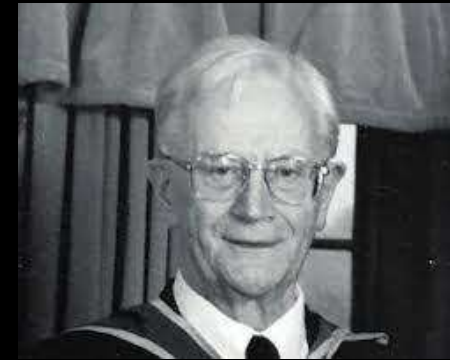
- Mode 2 ISWs
- Offshore wind
- Floating/submerged bridges
- Under water communications





# A bit about me

- BSc in Mathematics with European Studies: Durham University, 1995-1999
- Industrial Engineer: Viasystems Ltd, South Shields. 1999-2000
- PhD in Applied Mathematics: Durham University, 2000-2003
- Postdoctoral Research Assistant: University of Dundee, 2003-2006
- Lecturer/Senior Lecturer: University of St Andrews, 2006-2018
- Senior Lecturer: Newcastle University, 2018-present



Dr Vernon Armitage



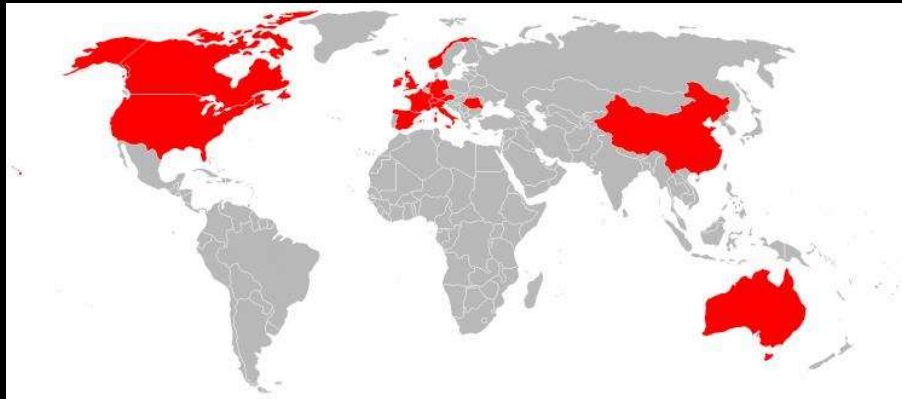
Prof Brian Straughan



Prof Peter. A. Davies

# Why have I stayed in academia ?

- World Expert in the field - highly specialised & original work.
- Opportunity to travel the world - meet people, discuss & present work, share your knowledge.



- Variety - Research, Teaching, Administration.
- Flexibility - Own boss. Hybrid working.
- Reflections - The times they are a changing, thanks to Piscopia and other initiatives.